

A model experiment of the quasi-biennial oscillation

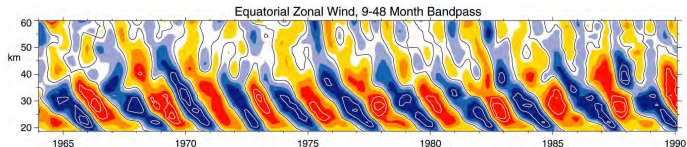
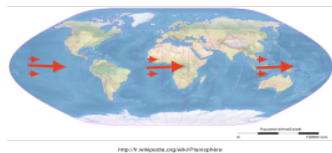
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Physics at the equator, ENS Lyon, October, 16th 2019

The atmospheric quasi-biennial oscillation (QBO)



Time-height section. Red : eastward wind. Blue : westward wind.

Baldwin et al. 2001

- Zonal (azimuthal) wind in the equatorial stratosphere (16-50 km).
- Amplitude $20 \text{ m} \cdot \text{s}^{-1}$
- Period about 28 months: not linked to any astrophysical forcing.
- Internal gravity waves/mean flow interaction. $T_{\text{wave}} < 2\text{days}$
- Similar winds on Jupiter and Saturn.

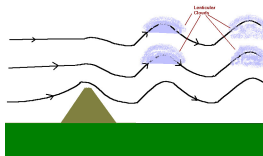
Internal gravity waves : geophysics

- Internal gravity waves propagate in stably **stratified fluids** subjected to **gravity**
- Inertia of the fluid. Restoring mechanism: gravity
- Ubiquitous in geophysical and astrophysical flows
- Lenticular clouds



i.pinimg.com/

- Mixing in the ocean



<http://www.prh.noaa.gov/hnl/pages/events/lenticular/>

Internal gravity waves

- Minimal model: Euler equation, mass conservation, incompressibility:

$$\rho \left(\frac{\partial \mathbf{v}}{\partial t} + (\mathbf{v} \cdot \nabla) \mathbf{v} \right) = -\nabla p + \rho \mathbf{g}$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{v}) = 0$$

$$\nabla \cdot (\mathbf{v}) = 0$$

Constant Brunt-Väisälä frequency:

$$N^2 = -\frac{d\rho_0}{dz} \frac{g}{\rho_0}$$

- Linearisation of the equations
Plane waves solutions
 $\mathbf{v}_0 = \mathbf{v} \exp(i(\mathbf{k} \cdot \mathbf{x} - \omega t))$

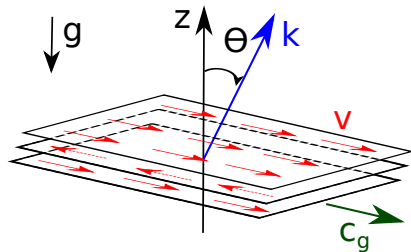
- Dispersion relation:

$$\omega^2 = N^2 \frac{k_x^2 + k_y^2}{k_x^2 + k_y^2 + k_z^2}$$

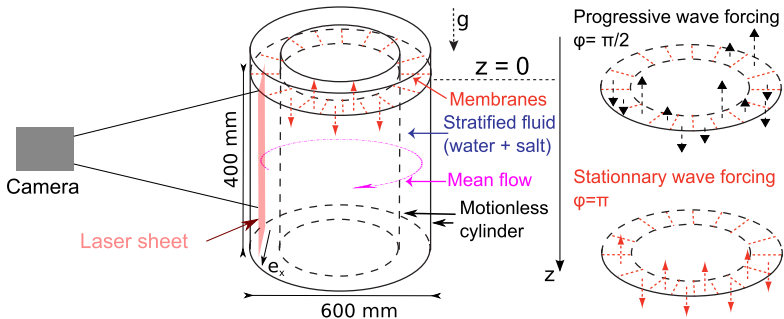
$$\omega = N |\sin(\theta)|$$

Anisotropic and dispersive waves

- From incompressibility: $\mathbf{v} \cdot \mathbf{k} = 0$
Group velocity direction: $\mathbf{c}_g \cdot \mathbf{k} = 0$



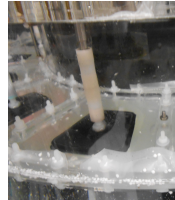
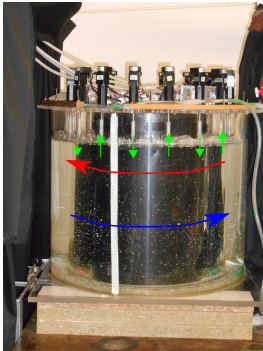
Experimental setup (1)



Plumb and McEwan (1978) , Otake et al. (1998)

- Väisälä-Brunt angular frequency $N = \sqrt{-\frac{d\rho_0}{dz} \frac{g}{\rho_0}} \sim 1.5 - 2.2 \text{ rad} \cdot \text{s}^{-1}$
- Forcing period $T \sim 15 - 38 \text{ s}$.
- Forcing amplitude: $M \leq 15 \text{ mm}$.
- Visualisation : particles, laser sheet. Particle image velocimetry (PIV)

Experimental setup (2)

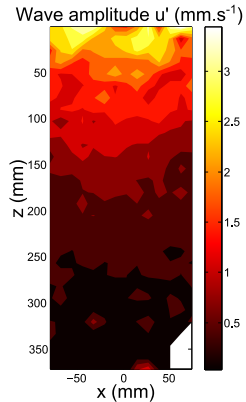
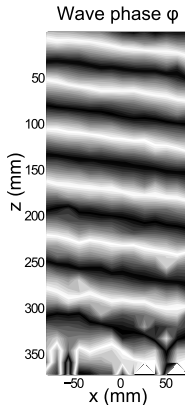
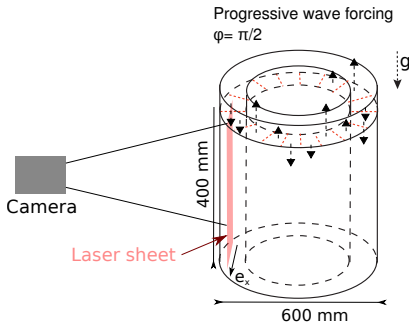


- Each membrane is controlled individually
- Continuous injection of fluid at the bottom and sucking of the fluid at the top (very small velocity $< 10 \mu\text{m} \cdot \text{s}^{-1}$): linear stratification despite of the mixing by the membrane motion.

Wave

Progressive wave forcing: 1 wave

Fit in each point of the horizontal velocity: $u(t) = u' \sin(\omega t + \varphi) + \bar{u}$
 with u' : wave amplitude, φ : initial phase, \bar{u} : mean flow

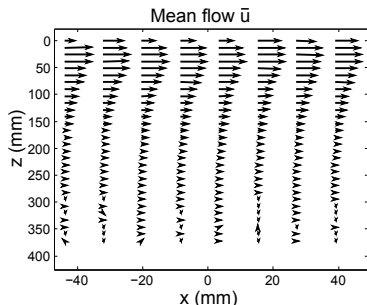


Forcing amplitude: $M = 8$ mm, $T = 28$ s, $N = 1.5$ rad · s⁻¹, steady state.

⇒ The component oscillating at ω has a wave structure.

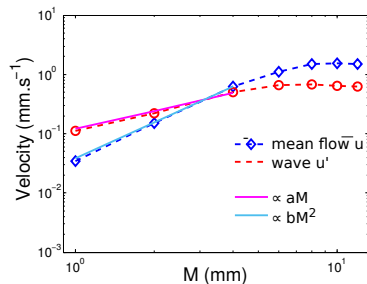
Mean flow

Fit in each point of the horizontal velocity: $u(t) = u' \sin(\omega t + \varphi) + \bar{u}$
with u' : wave amplitude, φ : initial phase, \bar{u} : mean flow



Forcing amplitude: $M = 8$ mm, $T = 28$ s,
 $N = 1.5 \text{ rad} \cdot \text{s}^{-1}$ steady state.

Analysis using Fourier Transform.



Forcing amplitude: $T = 38$ s,
 $N = 1.5 \text{ rad} \cdot \text{s}^{-1}$, $z = 64$ mm, steady state.

⇒ Always a mean flow.

⇒ The mean flow is homogeneous in the x direction (2D).

⇒ Mean flow proportional to the square of the wave at low M .

Variation with forcing amplitude M

- Navier-Stokes equation (3D):

$$\rho \left(\frac{\partial \mathbf{v}}{\partial t} + (\mathbf{v} \cdot \nabla) \mathbf{v} \right) = -\nabla p + \eta \Delta \mathbf{v} + \rho \mathbf{g}$$

- Velocity: $\mathbf{u} = (u' + \bar{u})\mathbf{e}_x + (v')\mathbf{e}_z$
- Model from Navier-Stokes:

$$\frac{\partial \bar{u}}{\partial t} = -\frac{\partial F}{\partial z} + \nu \frac{\partial^2 \bar{u}}{\partial z^2} - \gamma \bar{u}$$

with ν : kinematic viscosity

γ : wall friction

$$F = \overline{u'v'}$$

with v' the vertical wave component

- Prediction of the model: in steady state, without feedback of the mean flow on the wave: $\bar{u} \propto (u')^2$.
In agreement with experimental data.

Generation of an Eulerian mean flow by a wave



Acoustic streaming, W. Dridi et al.

Examples of mean flows induced by waves:

- Motion of liquid metals in electromagnetic pumps
- Eckhart acoustic streaming

- Decomposition of the velocity:

$$\mathbf{v} = \underbrace{\bar{\mathbf{V}}}_{\text{mean flow}} + \underbrace{\mathbf{v}'}_{\text{wave}} \quad (1)$$

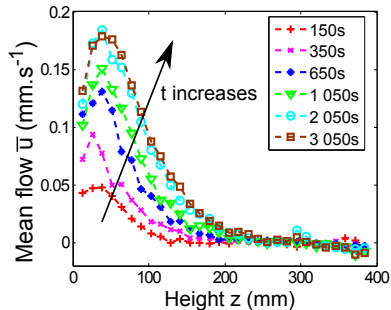
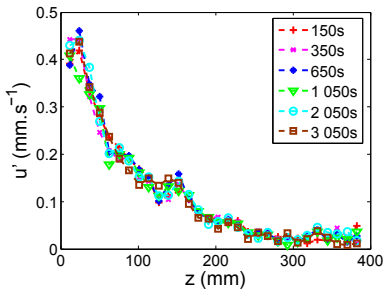
- Momentum equation (Navier-Stokes):

$$\rho \frac{d\bar{V}_i}{dt} = -\partial_i p + \eta \Delta V_i - \frac{\partial(\rho \overline{v'_i v'_j})}{\partial x_j} \quad (2)$$

- To generate a mean flow using waves in a steady regime, the waves must vary spatially (due to dissipation for example).

Linear regime

Linear regime, $M = 2$ mm



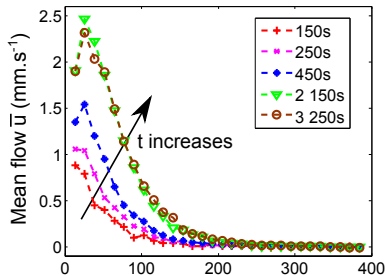
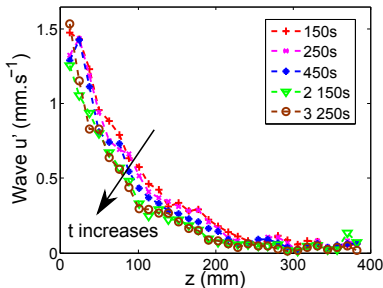
Forcing amplitude: $T = 38$ s, $M = 2$ mm, $N = 1.5$ rad · s⁻¹.

Constant wave amplitude u' : no feedback .

Mean flow increases and reaches a steady state in a timescale much longer than T .

Non-linear regime

Non-linear regime, $M = 8$ mm



Forcing amplitude: $T = 38$ s, $M = 8$ mm, $N = 1.5$ $\text{rad}\cdot\text{s}^{-1}$.

The wave amplitude u' is not constant : **feedback** of the mean flow on the wave.
Mean flow increases and reaches a steady state.

Feedback

- Model from Navier-Stokes, mass conservation, in Boussinesq and WKB approximations (from Plumb and McEwan 1978):

$$F(\xi) = F(0) \times \exp\left(-\int_0^\xi \left[\frac{1}{(1-\bar{U})^4}\right] d\xi\right)$$

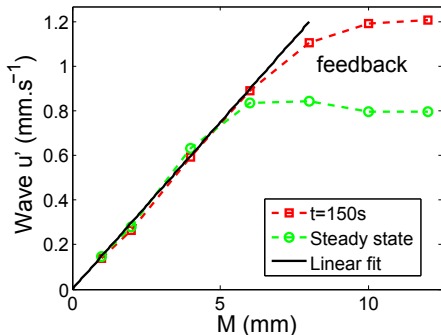
$$F = \overline{u'v'}$$

$$c = \frac{\omega}{k_x} \quad \text{horizontal phase velocity}$$

$$d = \frac{k_x c^4}{N^3 \nu} \quad \text{vertical dissipation length}$$

$$\xi = \frac{z}{d}$$

$$\bar{U} = \frac{\bar{u}}{c}$$



Forcing amplitude: $T = 38$ s, $N = 1.5$ rad · s⁻¹,
 $z = 64$ mm.

⇒ Negative feedback of mean flow on wave

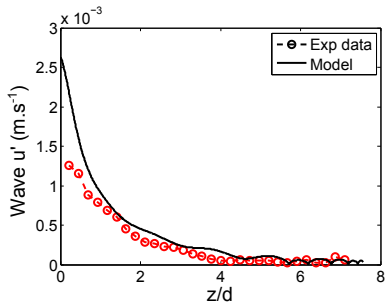
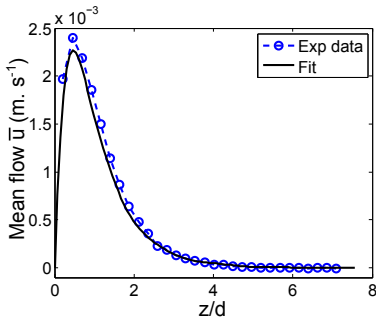
⇒ Feedback increases the momentum transfer from wave to mean flow at $z = 0$

Comparison with the model

$$\frac{\partial \bar{U}}{\partial t} = -d \frac{\partial F}{\partial \xi} + \nu d^2 \frac{\partial^2 \bar{U}}{\partial \xi^2} - \gamma \bar{U}$$

$$F(\xi) = F(0) \times \exp\left(-\int_0^\xi \left[\frac{1}{(1-\bar{U})^4}\right] d\xi\right)$$

Comparison with the model. One single fitting parameter $F(0)$.



Forcing amplitude: $T = 38$ s, $M = 8$ mm, $N = 1.5$ rad · s⁻¹, steady state.

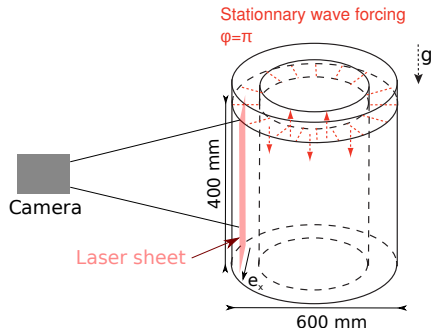
Good agreement between the model and experimental data

Progressive wave forcing: conclusion

- Dissipation of the wave due to viscosity
- Generation of a mean flow
- Negative feedback of the mean flow on the wave
- 1D model sufficient

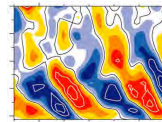
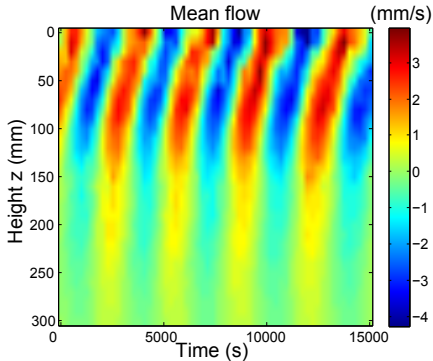
Stationnary wave forcing

Stationnary wave forcing: 2 waves



Mean flow

Stationnary wave forcing: 2 waves



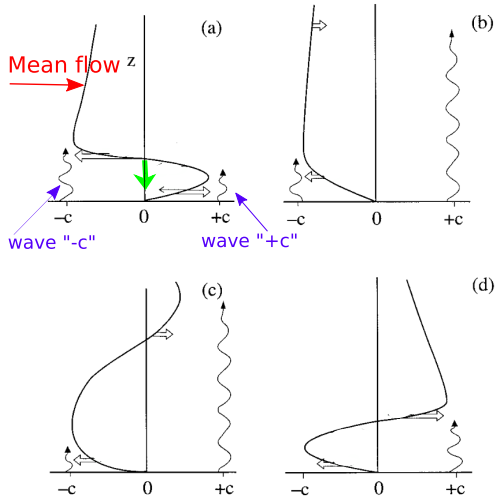
Oscillation in the
atmosphere

Forcing amplitude: $M = 13.5 \text{ mm}$, $T = 17 \text{ s}$,
 $N = 1.5 \text{ rad} \cdot \text{s}^{-1}$, $T_{\text{QBO}} = 3200 \text{ s}$

- ⇒ No mean flow below threshold forcing amplitude.
- ⇒ Observation of an oscillation similar to the one of the atmosphere

Basic mechanisms of the QBO

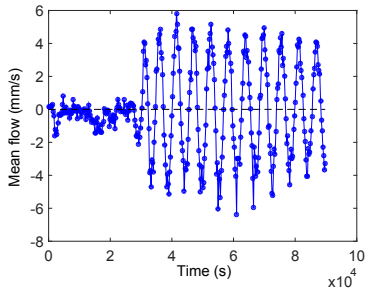
- 2 internal gravity waves with opposite phase velocity in the azimuthal direction
- Dissipation of the waves.
- The transfer of momentum from the wave to the mean flow is more efficient if c and the mean flow V are of the same sign (positive feedback).
- Momentum transferred from a wave to the mean flow can not be transferred above.
- Dissipation of the large mean flow gradient.



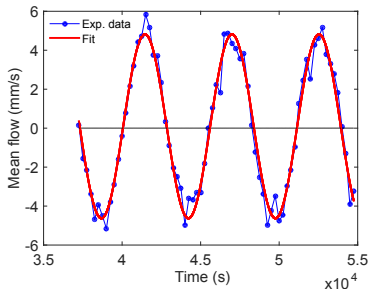
Adapted from Baldwin et al 2001

Mean flow at a given height

Height: $z = 88$ mm



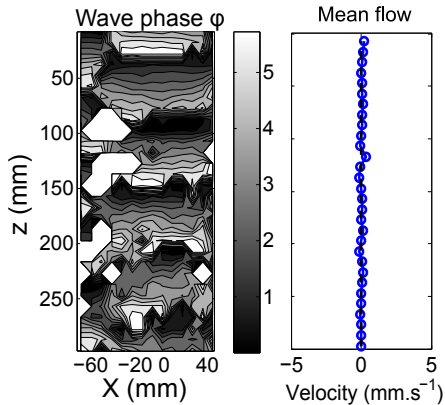
Forcing amplitude: $M = 13.5$ mm, $T = 15$ s,
 $N = 1.5$ rad \cdot s $^{-1}$, NaCl



Cosine fit : $V = A \cos(\omega t + \Phi) + V_n$

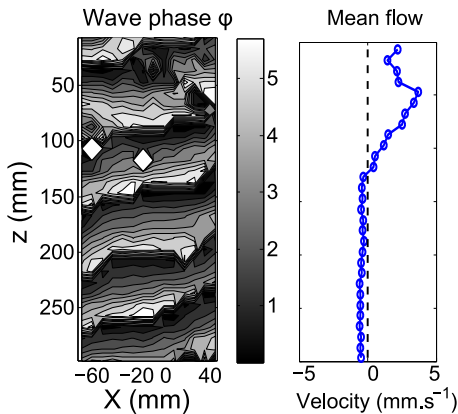
The mean flow may appear after a very long time.
Sine oscillation of the mean flow.

Wave without mean flow



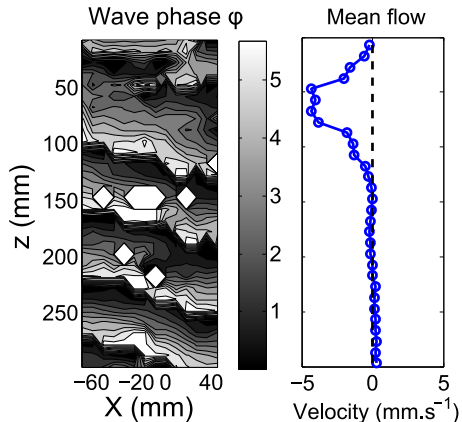
Initial state, no mean flow.

Wave and mean flow



Permanent state, time t_1

⇒ Feedback of the mean flow on the waves



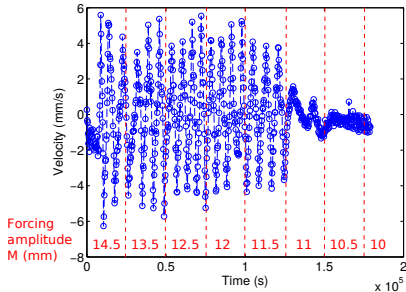
Permanent state, time $t_1 + 2\,500\text{ s}$

Influence of the stratification

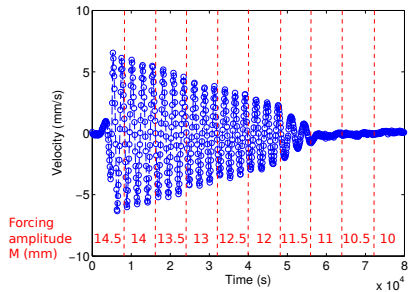
$T = 15$ s, $z = 30$ mm

Mean flow

$N = 1.5 \text{ rad} \cdot \text{s}^{-1}$, NaCl



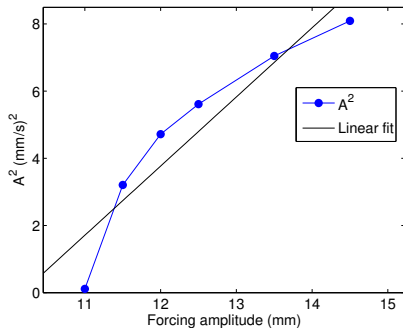
$N = 2.15 \text{ rad} \cdot \text{s}^{-1}$, MgCl_2



Bifurcation

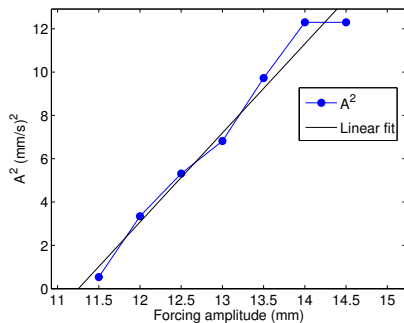
$T = 15$ s, A : vertical average of mean flow amplitude

$N = 1.5 \text{ rad} \cdot \text{s}^{-1}$, NaCl



Subcritical bifurcation

$N = 2.15 \text{ rad} \cdot \text{s}^{-1}$, MgCl_2



Supercritical bifurcation

⇒ Nature of the bifurcation depends on the parameters

Yoden and Holton (1988)

Dimensionless model

$$\frac{\partial U}{\partial T} = -\frac{\partial D}{\partial \xi} + \Lambda_1 \frac{\partial^2 U}{\partial \xi^2} - \Lambda_2 U \quad (3)$$

Forcing with two counter-propagating waves:

$$D(\xi) = \exp\left(-\int_0^\xi \frac{1}{(1-U)^4} d\xi\right) - \exp\left(-\int_0^\xi \frac{1}{(1+U)^4} d\xi\right) \quad (4)$$

with the scales:

- velocity: $c = \omega/k_x$
- length: dissipation length $d = (k_x c^4)/(N^3 \nu)$
- time: $cd/(F(0))$

and the parameters:

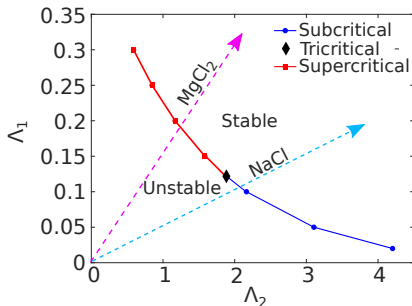
- $\Lambda_1 = \nu c/(F(0)d)$
- $\Lambda_2 = \gamma cd/(F(0))$

This model can be solved semi-analytically close to the threshold.

Threshold

$$\frac{\partial U}{\partial T} = -\frac{\partial D}{\partial \xi} + \Lambda_1 \frac{\partial^2 U}{\partial \xi^2} - \Lambda_2 U$$

- $\Lambda_1 = \nu c / (F(0)d)$
- $\Lambda_2 = \gamma cd / (F(0))$



Yoden and Holton 1988

⇒ Hopf bifurcation

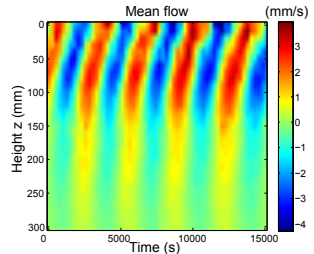
⇒ Nature of the bifurcation depends on the form of the dissipation

Conclusion

- A damped progressive internal wave induces a mean flow
- Feedback of the mean flow on the wave at high forcing amplitude
- Threshold for mean flow in symmetric forcing (2 counter-propagating waves)
- Observation of mean flow reversal.
- Nature of the bifurcation is supercritical or subcritical, depending on dissipation

Perspectives:

- Link with atmospheric QBO: bifurcation in general circulation models (GCM) and in atmosphere.



B. Semin, N. Garroum, F. Pétrélis, S. Fauve, "Nonlinear saturation of the large scale flow in a laboratory model of the quasibiennial oscillation", *Phys. Rev. Lett.* **121**, 134502 (2018)

B. Semin, G. Facchini, F. Pétrélis, S. Fauve, "Generation of a mean flow by an internal wave", *Phys. Fluids*, **28**, 096601 (2016).